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A. Šoštarić, S. Stanišić Stojić, G. Vuković, Z. Mijić, A. Stojić, I. Gržetić

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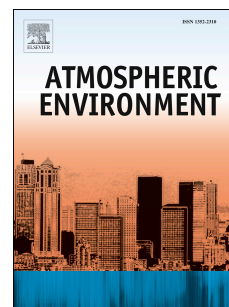
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## 1 Rainwater capacities for BTEX scavenging from ambient air

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3 A. Šoštarić<sup>1\*</sup>, S. Stanišić Stojić<sup>2</sup>, G. Vuković<sup>3</sup>, Z. Mijić<sup>3</sup>, A. Stojić<sup>3</sup> and I. Gržetić<sup>4</sup>

4

5 *1 Institute of Public Health Belgrade, Bulevar Despota Stefana 54a, 11000 Belgrade, Serbia*6 *2 Faculty of Physical Chemistry, University of Belgrade, Studentski trg 12-16, 11000 Belgrade, Serbia*7 *3 Institute of Physics Belgrade, University of Belgrade, Pregrevica 118, 11080 Belgrade, Serbia*8 *4 Faculty of Chemistry, University of Belgrade, Studentski trg 12-16, 11000 Belgrade, Serbia*9 \*Corresponding author: [andrej.sostaric@zdravlje.org.rs](mailto:andrej.sostaric@zdravlje.org.rs); Phone: +381 64 13 94 185; Fax: +381 11 32 35 080;  
10 Bulevar Despota Stefana 54a, 11000 Belgrade, Serbia.

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## 37 Abstract

38 The contribution of atmospheric precipitation to volatile organic compound (VOC) removal from the  
39 atmosphere remains a matter of scientific debate. The aim of this study was to examine the potential  
40 of rainwater for benzene, toluene, ethylbenzene and xylene (BTEX) scavenging from ambient air. To  
41 that end, air and rainwater samples were collected simultaneously during several rain events that  
42 occurred over two distinct time periods in the summer and autumn of 2015. BTEX concentrations in  
43 the gaseous and aqueous phases were determined using proton transfer reaction mass  
44 spectrometry. The results reveal that the registered amounts of BTEX in rainwater samples were  
45 higher than those predicted by Henry's law. Additional analysis, including physico-chemical  
46 characterization and source apportionment, was performed and a possible mechanism underlying  
47 the BTEX adsorption to the aqueous phase was considered and discussed herein. Finally, regression  
48 multivariate methods (MVA) were successfully applied (with relative errors from 20%) to examine  
49 the functional dependency of BTEX enrichment factor on gaseous concentrations, physico-chemical  
50 properties of rainwater and meteorological parameters.

51 **Keywords:** BTEX, wet deposition, rain, PTR-MS, multivariate methods, Unmix.

52

## 53 1. Introduction

54 Benzene, toluene, ethylbenzene and the three xylene isomers, frequently referred to as BTEX,  
55 constitute a group of aromatic hydrocarbon species of particular environmental interest, commonly  
56 associated with the petrochemical industry and incomplete fossil fuel oxidation (Stojić et al. 2015a;  
57 Stojić et al., 2015b). Besides being important photochemical precursors for tropospheric ozone and  
58 secondary organic aerosols (SOA) (Chatani et al., 2015), these hazardous air pollutants cause chronic  
59 toxicity even in small concentrations (Stojić et al., 2015c). According to the IARC data, benzene is  
60 recognized as a significant public health threat and classified as group I carcinogen, ethylbenzene is a  
61 suspected IIB carcinogen, while both toluene and the xylene isomers belong to group III neurotoxins  
62 (WHO, 1986, 1993, 1997; Durmusoglu et al., 2010).

63 In the atmosphere, volatile species are distributed between the gaseous, aqueous and particle phase  
64 (Matsumoto et al., 2010). In their biogeochemical cycle, it is believed that the role of atmospheric  
65 water is quite prominent, but this issue is still subject to continuous scientific debate (McNeill et al.,  
66 2012). The concentrations of BTEX in various forms of atmospheric water depend on various factors  
67 including their ambient gas mixing ratios, water solubility and Henry's law constant, frequency and  
68 intensity of precipitation events (Balla et al., 2014), gas-water surface interactions (Raja and Valsaraj,  
69 2004), content and concentrations of other species in atmospheric water (Okochi et al., 2005; Sato et  
70 al., 2006; Allou et al., 2011), as well as the origin of air masses (Mullaugh et al., 2015). Previous  
71 studies, which primarily focused on wet deposition of BTEX and their partition between gaseous and  
72 aqueous phases, were relatively scarce and provided contradictory conclusions.

73 In the study aimed at investigating the capacity of rainwater for wet scavenging of BTEX, Okochi et al.  
74 (2004) reported that the concentrations of species detected in rain samples were higher than those  
75 predicted by Henry's law, and concluded that atmospheric precipitation might play significantly  
76 greater role in removing BTEX from ambient air than previously thought. Thereby, the observed  
77 supersaturation was assumed to be associated with the presence of surface-active agents in rain  
78 droplets, whereas the rainfall intensity appeared to be of negligible importance. Accordingly, our  
79 previous study confirmed a significant enrichment of BTEX in the aqueous phase in a dynamic  
80 equilibrium system designed to resemble the interactions between the gaseous and water phase  
81 during rainfall (Šoštarić et al., 2016). Conversely, recent findings of Mullaugh et al. (2015) indicate  
82 that BTEX were not efficiently scavenged from the atmosphere by wet deposition processes.  
83 Furthermore, the authors concluded that light-mediated reactions with OH· or nitrogen radicals  
84 remain the major atmospheric sink for BTEX. Nonetheless, it should be noted that this research was

85 not based on the ambient air measurements, but it mainly relied on the previously published BTEX  
86 data from similar locations.

87 In order to better understand the fate of volatile species in atmospheric, terrestrial and aquatic  
88 systems, the present study examines the contribution of rainwater to wet scavenging of atmospheric  
89 BTEX, as well as the mechanisms related to their air-water distribution transfer.

90

## 91 **2. Materials and methods**

92 A total of 53 sample pairs of air and rainwater samples were collected simultaneously during several  
93 rain events that occurred over two distinct time periods in the summer and autumn season of 2015.  
94 The sampling was performed at the Institute of Physics (Belgrade, Serbia; 44°49' N, 20°28' E), located  
95 in the vicinity of the Danube river, in the suburban residential area, with a number of local fireboxes  
96 active during the heating season.

97 Rainwater sampling was performed using a custom-built precipitation collector with the effective  
98 sampling area of 9 m<sup>2</sup>. The steep collecting panels (45°) were designed to reduce rainfall retention  
99 time and minimize possible BTEX volatilization. Such large sampling area enabled collecting a vast  
100 number of samples per each rain event. The panels were thoroughly rinsed with 18 MΩ ultrapure  
101 water (ELGA PURELAB maxima system) prior to each sampling campaign, and the rinsing water was  
102 collected and analyzed as a field blank control sample. No target compounds were detected in the  
103 field blank control samples. The samples were collected and stored directly into brown glass bottles  
104 of 1300 mL. All bottles were washed with detergent, thoroughly rinsed with ultrapure water and  
105 dried in an oven for two hours at 105°C to remove any trace of contamination. During the sampling,  
106 the bottles were filled to the top to avoid headspace, and the sampling duration and sample  
107 temperature were recorded. Since sampling equipment enables collection of large volumes of  
108 rainwater within a short period, the last sample in each sampling campaign was collected in the  
109 bottle of 2,600 mL and was split into two standard aliquots. The first aliquot was analyzed  
110 immediately, whereas the other one was examined after all other samples to determine whether the  
111 BTEX levels changed over time. No difference could be observed in the obtained quantity of double  
112 samples (Table S1, Supplementary material).

113 BTEX concentrations in both gas and water phases were measured using proton transfer reaction  
114 mass spectrometer (Standard PTR-quad-MS, Ionicon Analytik, GmbH, Austria), whose detailed  
115 description is given elsewhere (Lindinger et al., 1998). Since PTR-quad-MS is not capable of  
116 distinguishing isobaric ions, the signal detected at m/z 107 referred to C<sub>8</sub> aromatic hydrocarbons,  
117 ethylbenzene, o-, m-, and p-xylene. Signals detected at m/z 79 and m/z 93 referred to benzene and  
118 toluene, respectively (Warneke et al., 2003).

119 The air samples were collected as a side flow from a 1/8-inch teflon tube sampling line through which  
120 ambient air was drawn at the flow rate of about 50 L min<sup>-1</sup> to ensure short residence. The sample  
121 inlet was located 6 m above ground level with a sampling angle of 360°. Drift tube parameters  
122 included: pressure, ranging from 2.04 to 2.14 mbar; temperature, 60 °C; voltage, 600 V; E/N  
123 parameter, 145 Td providing reaction time of 90 μs. The count rate of H<sub>3</sub>O<sup>+</sup>H<sub>2</sub>O was 3 to 8% of the  
124 9.2·10<sup>6</sup> counts s<sup>-1</sup> count rate of primary H<sub>3</sub>O<sup>+</sup> ions. PTR-MS calibration was performed before and  
125 after each sampling campaign using an external standard five-point calibration, ranging from  
126 0-26 ppbV, 0-25 ppbV and 0-80 ppbV for B, T and EX, respectively. For this purpose, 2.5 ppmV  
127 mixture of BTEX (BTEX in nitrogen, Messer Group GmbH) was diluted with high-purity synthetic air  
128 (CH free, Messer Group GmbH) by means of HORIBA ASGU 370-P system.

129 Determination of BTEX concentrations in rainwater was performed immediately after each sampling  
130 campaign. A liter of each unfiltered rainwater sample was transferred to the gas washing bottle  
131 (GWB) and purged out with synthetic air at a flow rate of 1 L min<sup>-1</sup>. Rainwater filtration was avoided  
132 due to potential adsorption of species on the filter. The GWB output was connected with PTR-MS

133 inlet via T-piece, and further analytical procedure, calibration and data processing were conducted as  
 134 described in Šoštarić et al. (2016). In brief, PTR-MS signal obtained during exsufflation was subject to  
 135 baseline fitting. The exsufflation time was determined for each sample as the interval required for  
 136 equilibrium to be achieved ( $t_{eq}$ ). The obtained exsufflation time was used for determining the  
 137 amounts of target compounds retained in the analyzed rainwater samples. The aqueous  
 138 concentrations of analyzed species were calculated by multiplying the obtained amounts by the  
 139 conversion factor (3.25; 3.83 and 4.41 for B, T and EX, respectively). The detection limits (DL) in  
 140 rainwater were determined using HC free air and calculated as 10 nM, 10 nM and 20 nM for B, T and  
 141 EX, respectively. The remaining portion of each rainwater sample was transferred to a 300-mL bottle  
 142 and stored at 4°C until further analysis, which included determination of the major inorganic anions  
 143 ( $F^-$ ,  $Cl^-$ ,  $SO_4^{2-}$ ,  $NO_2^-$  and  $NO_3^-$ ), dissolved cations ( $Na^+$ ,  $NH_4^+$ ,  $K^+$ ,  $Ca^{2+}$  and  $Mg^{2+}$ ), total organic carbon,  
 144 electrical conductivity, UV extinction, turbidity and pH, in accordance with the standard methods (US  
 145 EPA 300.1:1993, EN ISO 14911:1998, ISO 8245:1999, EN 27888:1993, SMEWW 19th method 5910 B,  
 146 US EPA 180.1:1993, EN ISO 10523:2008, respectively). More details of the methods and equipment  
 147 applied for physico-chemical analysis conducted on rainwater samples are presented in  
 148 Supplementary material.

149 In order to determine the extent to which Henry's law constant ( $K_H$ ) describes BTEX distribution  
 150 between the gaseous and aqueous phase, distribution coefficients ( $D_{OBS}$ ) were calculated for each  
 151 sample pair and each species, as the ratio of the corresponding experimentally derived rainwater  
 152 concentrations in nM ( $C_R$ ) and ambient gas phase mixing ratios in ppbV ( $p_g$ ):

$$153 \quad D_{OBS} = C_R / p_g \quad (M \text{ atm}^{-1}) \quad (1)$$

154 Furthermore, the enrichment factors (EF) were calculated as the ratio of  $D_{OBS}$  and  $K_H$ .

155 Considering the  $K_H$  temperature dependence, EF were calculated using temperature corrected  $K_H T$   
 156 for each rain sample by means of the following equation (Sander, 2015):

$$157 \quad K_H T = K_H(298.15) \exp \left\{ \frac{-\Delta H}{R} \left( \frac{1}{298.15} - \frac{1}{T} \right) \right\} \quad (2)$$

158 where  $K_H$  is the Henry's law constant at 298.15 K for pure water,  $\Delta H$  is the enthalpy change of air-  
 159 water transfer,  $T$  is the rainwater temperature, and  $R$  is the universal gas constant ( $8.314 \text{ J K}^{-1} \text{ mol}^{-1}$ ).  
 160 Furthermore, to assess the representativeness of ground level conditions for the atmospheric  
 161 conditions during rainfall,  $K_H T$  and EF altitude profiles were calculated using the temperature profiles  
 162 obtained from GDAS1 (Global Data Assimilation System, 2015), by replacing the rainwater  
 163 temperature value by the temperature at the corresponding altitude.

164 Meteorological parameters during rain events (precipitation (accumulated rainfall, rain current and  
 165 peak intensity, and the duration of a rain event), wind speed and direction, pressure, humidity and  
 166 temperature) were measured by Vaisala Weather Transmitter WXT530 Series. Cloud information,  
 167 including cloud height and type, was obtained from the airport "Nikola Tesla", Belgrade, ICAO code  
 168 LYBE, located 8.9 km SSW from the sampling site.

169 The relationships between enrichment factors (EF), physico-chemical characteristics and wind  
 170 characteristics (wind speed and direction) were examined using the bivariate polar plot analyses  
 171 (Carslaw and Beevers, 2013) implemented in the Openair package (Carslaw and Ropkins, 2012)  
 172 within the statistical software environment R (Team, 2012).

173 The neutralization factors (NF) were calculated based on the study of Moreda-Piñeiro et al. (2014),  
 174 Tiwari et al. (2016) and references therein in order to examine the potential of the cations to balance  
 175 the rainwater acidic components:

$$176 \quad [NF_{Ca^{2+}}] = \frac{[nssCa^{2+}]}{NO_3^- + [nssSO_4^{2-}]} \quad (3)$$

$$177 \quad [NF_{Mg^{2+}}] = \frac{[nssMg^{2+}]}{NO_3^- + [nssSO_4^{2-}]} \quad (4)$$

$$178 \quad [NF_{NH_4^+}] = \frac{[NH_4^+]}{NO_3^- + [nssSO_4^{2-}]} \quad (5)$$

$$179 \quad [NF_{K^+}] = \frac{[nssK^+]}{NO_3^- + [nssSO_4^{2-}]} \quad (6)$$

180 To calculate the non-sea salt fraction of any particular ion (*nss*), we assumed that all Na originated  
181 from marine sources, and used it as a referent element. The *nss* contribution is given as:

$$182 \quad [nss - X]_i = [X]_i - [Na^+]_i \left[ \frac{[X]}{[Na^+]} \right]_{sea\ salt} \quad (7)$$

183 where  $[nss - X]_i$  is the *nss* concentration of the selected ion in the sample *i*,  $[X]_i$  is the total  
184 concentration of the ion X measured in the rainwater sample *i*,  $[Na^+]_i$  is the total concentration of  $Na^+$   
185 measured in the rainwater sample, and  $\left[ \frac{[X]}{[Na^+]} \right]_{seasalt}$  is the reference ratio determined in the  
186 seawater.

187 Potential remote source regions that might affect the observed BTEX mixing ratios were identified  
188 using HYSPLIT-derived 72-h back trajectories (Draxler and Rolph, 2014). The trajectories were  
189 computed for each hour UTC a day before and during each rain event, above the sampling location at  
190 the half of the planetary boundary layer height calculated from GDAS1 using Meteoinfo (Wang,  
191 2014), as described in Stojić et al. (2016) and Stojić and Stanišić Stojić (2017).

192 Rainwater source apportionment was performed using Unmix (USEPA, 2007). The maximum number  
193 of species selected as input variables was chosen using the trial and error with the overall aim of  
194 yielding the most physically meaningful results. For concentrations below the DL, a value equal to the  
195 half of the DL was used.

196 Guided regularized random forest (GRRF) was applied (Deng and Runger, 2013) for the selection of  
197 features that are most relevant for EF. Random forest (RF) consists of a number of decision trees  
198 which every node represents a condition on a single variable designed to split the dataset in two  
199 parts so that similar response values end up in the same set. Variable importance measures how  
200 much each variable decreases the weighted impurity across all trees, a measure based on which the  
201 optimal condition is chosen. GRFF uses the importance scores from a preliminary RF to guide the  
202 feature selection of regularized random forest (RRF), and has several advantages as follows: it is  
203 more robust and computationally efficient than RRF, varSelRF and LASSO logistic regression; it can  
204 select compact feature subsets moderating the curse of dimensionality; it avoids the effort to analyze  
205 irrelevant or redundant features; and it has competitive accuracy performance. Variable importance  
206 presented herein was obtained by calculating the average value of 2000 GRRF runs. The appropriate  
207 number of trees was determined to assure out-of-bag error convergence. Method performance was  
208 tested by 100 times replicated 10-fold cross validation.

209 To analyze the relationship between EF and features that are considered most relevant for EF  
210 prediction, the following 24 regression MVA methods implemented in Weka 3.8 (Hall et al., 2009)  
211 were applied: Alternating Model Tree, Conjunctive Rule, Decision Stump, Decision Table, Elastic Net,  
212 Gaussian Processes, IBk, IBkLG, Isotonic Regression, K\*, Least Median Squared, Linear Regression,  
213 Locally Weighted Learning, M5P, M5 Rules, Multilayer Perceptron, Pace Regression, Random Forest,  
214 Random Tree, Radial Base Function (RBF) Network, RBF Regressor, REP Tree, Simple Linear  
215 Regression and SMOreg Support Vector Machine. A brief description of the methods, including  
216 functions (neural network, support vector machine, etc.), clustering techniques, rules and trees, is  
217 provided in Supplementary material. Method performance was tested by 10 times replicated 10-fold  
218 cross-validation.

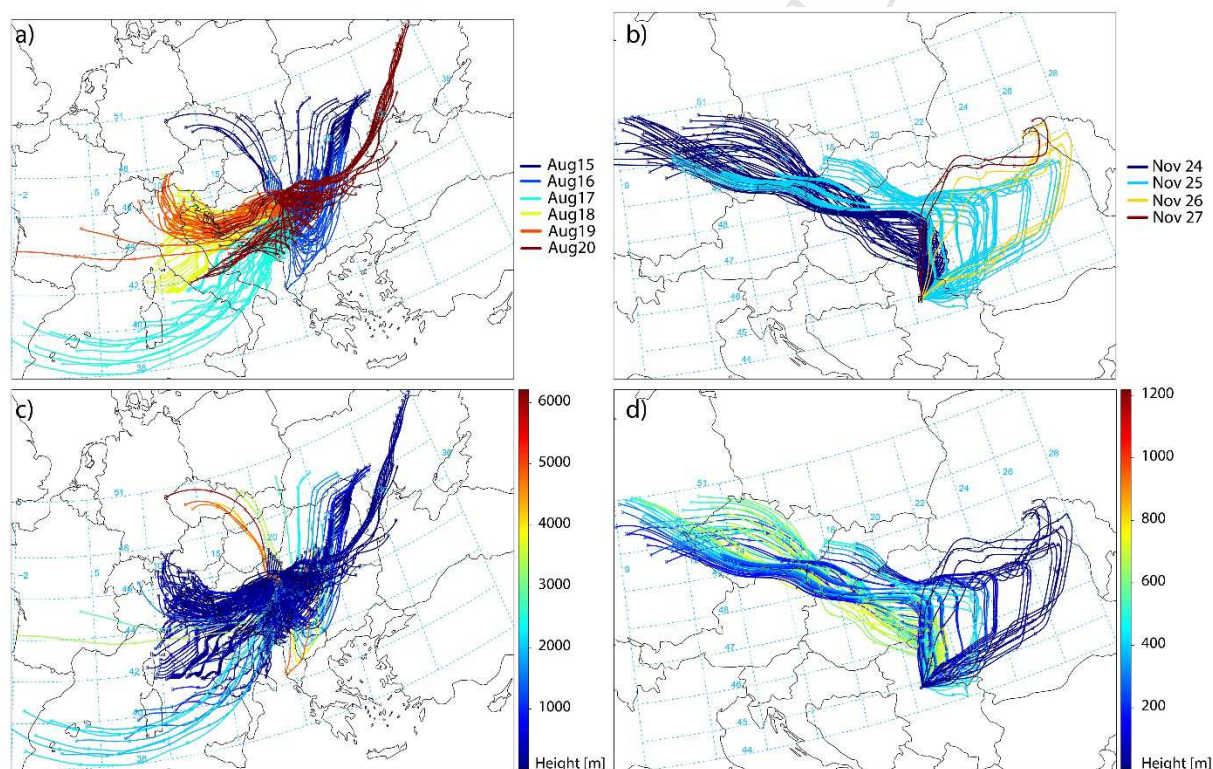
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### 220 3. Results and discussion

221 Light showers, with occasional thunderstorms, constituted a considerable part of summer rain  
 222 events. Scattered and broken clouds in the form of cumulonimbus or towering cumulus were  
 223 observed at the height of 400 to 1000 m. In the autumn campaign, the vast majority of rain events  
 224 were light and sporadically followed by mist. Scattered clouds were registered at levels below 300 m,  
 225 whereas broken clouds were observed from 100 to 900 m.

226 As expected, both aqueous and gaseous BTEX concentrations were higher during the cold part of the  
 227 year. Lower BTEX concentrations in summer can be attributed to intense photochemical removal and  
 228 washout effects associated with more sunny and rainy days (Lee et al., 2002), whereas higher  
 229 concentrations of BTEX in autumn can be associated with individual combustion fireboxes widely  
 230 spread in the vicinity of the sampling site. Furthermore, aqueous B concentrations in summer were  
 231 below DL.

232 As can be seen in Figure 1, each rain event was associated with air masses coming from different  
 233 source regions and heights. During summer rain events, air flows from all directions were followed by  
 234 significant variations in physico-chemical properties of rainwater, whereas N and NE air masses in  
 235 autumn were associated with a more uniform rainwater composition. Gaseous concentrations of  
 236 volatile species, particularly B, were relatively stable during rain events, which can be explained by  
 237 the fact that the sampling site was dominated by local BTEX sources.



238  
 239 Figure 1. Back trajectories a day before (August 15 and November 24) and during summer (a) and  
 240 autumn (b) rain events and corresponding trajectory heights (c, d).

### 241 3.1 Physico-chemical characteristics of rainwater

242 Basic statistics for all rainwater parameters and BTEX concentrations is given in Table S2,  
 243 Supplementary material. The average rainwater pH was 6.01, while the turbidity was below 10 NTU,  
 244 indicating that the samples contained moderate amounts of suspended particles from the  
 245 atmosphere. As illustrated in Figure S1, conductivity, as well as high concentrations of most ions  
 246 ( $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$ ,  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{K}^+$  and  $\text{Na}^+$ ) were influenced by the high-speed SW wind (20 to 30  $\text{m s}^{-1}$ ),  
 247 while only  $\text{NH}_4^+$  concentrations were increased with the wind of moderate speed (10  $\text{m s}^{-1}$ ) from NE  
 248 direction.

249 The rainwater pH varied from 3.70 to 8.20 with the volume weighted mean of 6.01, which is mainly  
 250 due to scavenging of alkaline species ( $\text{Ca}^{2+}$  and  $\text{SO}_4^{2-}$ ). The average pH value is also close to the 5-  
 251 year-mean (6.1) obtained as a part of the regular air quality monitoring in Belgrade. The contribution  
 252 of  $\text{SO}_4^{2-}$  to the rainwater acidity was confirmed by ( $\text{Cl}^- + \text{NO}_3^-$ ) and ( $\text{SO}_4^{2-}$ ) ratio below 1 (Tiwari et al.,  
 253 2016). As shown,  $\text{Ca}^{2+}$  was the dominant neutralization component, followed by  $\text{NH}_4^+$ ,  $\text{Mg}^{2+}$  and  $\text{K}^+$ ,  
 254 with mean of 77%, 14%, 7% and 2.3%, respectively.

255 As shown in Figure S2, significant correlations were observed as follows:  $> 0.80$  ( $\text{NO}_3^- - \text{SO}_4^{2-}$ ,  $\text{NH}_4^+ -$   
 256 aq. B,  $\text{NH}_4^+ - \text{gas. B}$ ,  $\text{NH}_4^+ - \text{EF}_B$ );  $0.70 - 0.80$  ( $\text{F}^- - \text{SO}_4^{2-}$ ,  $\text{F}^- - \text{NO}_3^-$ ,  $\text{F}^- - \text{aq. B}$ ,  $\text{F}^- - \text{gas. B}$ ,  $\text{F}^- - \text{EF}_B$ ,  $\text{K}^+ -$   
 257  $\text{Na}^+$ ); and  $0.60 - 0.70$  ( $\text{NO}_3^- - \text{Mg}^{2+}$ ,  $\text{SO}_4^{2-} - \text{Mg}^{2+}$ ,  $\text{SO}_4^{2-} - \text{EF}_{\text{EX}}$ ,  $\text{Na}^+ - \text{Mg}^{2+}$ ). Furthermore, high  
 258 correlations were noted between aq. and gas. BTEX concentrations ( $\geq 0.80$ ), as well as between aq.  
 259 EX and aq. B (0.67) and aq. T (0.67), suggesting that these species might share the common source.

260 As can be seen in Tables S3, S4 and S5, four factors were derived using Unmix. With high  
 261 contributions of volatile BTEX species (99%, 44.5% and 52.2%) and a relatively high share of UV  
 262 extinction (33.3%), the first factor was recognized as organic compounds in the gaseous form. The  
 263 second factor, characterized by the highest contribution of TOC (71.6%) and turbidity (79.7%),  
 264 represented the solid fraction dissolved in the atmospheric water. Significant shares of  $\text{K}^+$  (58.0%),  
 265  $\text{SO}_4^{2-}$  (42.2%) and  $\text{NO}_3^-$  (31.3%) were also apportioned to this factor, as a result of fossil fuel burning  
 266 and traffic exhaust (Rao et al., 2016; Tiwari et al., 2016). The high shares of crustal-related elements,  
 267  $\text{Na}^+$  (61.8%) and  $\text{Mg}^{2+}$  (61.4%), were apportioned to the third factor (Cao et al., 2008; Sapek, 2014).  
 268 Moderate to significant shares ( $> 30\%$ ) of all species except  $\text{Na}^+$ ,  $\text{K}^+$  and B were apportioned to the  
 269 fourth factor, being recognized as the aerosol fraction. Apart from gaseous oxides ( $\text{SO}_2$  and  $\text{NO}_2$ ),  
 270 which react with ozone and OH· radicals in the presence of  $\text{Ca}^{2+}$  and  $\text{Mg}^{2+}$  (Seinfeld and Pandis, 2006),  
 271 BTEX are susceptible to photo-oxidation that can also lead to SOA formation. BTEX reactions include  
 272 oxidation with ozone and OH·, but also with  $\text{NO}_x$  and  $\text{SO}_2$ , which results in multi-functional oxy  
 273 products that are further deposited onto the existing aerosol or initiate the formation of SOA by self-  
 274 nucleation. BTEX behave differently in the atmosphere due to differences in the methyl chain  
 275 substituent and the alkyl chain length. Benzene is considered being extremely stable compared to T  
 276 and EX (Słomińska et al., 2014), and it is less susceptible to the heterogeneous reactions and  
 277 formation of SOA in the atmosphere.

278 The contributions of the gaseous organic- and aerosol-related factors were mostly associated with N  
 279 wind of moderate speed ( $< 10 \text{ m s}^{-1}$ ), which clearly reflects their local origin, while the solid- and  
 280 crustal-related factors were associated with the air masses from SW direction and high wind speed  
 281 ( $20$  to  $30 \text{ m s}^{-1}$ ) (Figure S3). The contribution of the factor assigned to aerosols declined during rain  
 282 events due to wet deposition, while similar behavior was not observed for other factors.

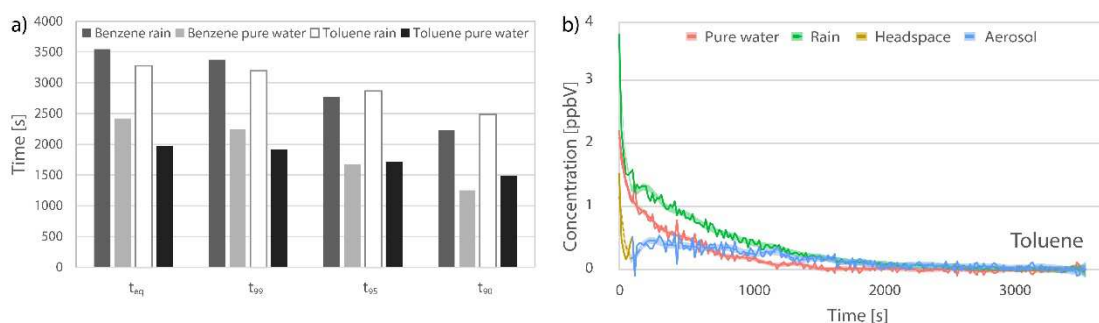
283 TEX air mixing ratios and rain concentrations decreased during the first two hours of the rainfall, but  
 284 tend to rise afterwards probably as a result of rainfall intensity decrease (Figure S4, Supplementary  
 285 material). The highest TEX enrichment, caused by air mixing ratio decrease, was detected during the  
 286 second hour. Typical washout effect was observed for source contributions related to the rainwater  
 287 aerosol and solid components, and was less pronounced for crustal factor. Unlike TEX, only a slight  
 288 decrease in benzene air mixing ratios was noticeable, which is reflected in a constant  $\text{EF}_B$  increase.

### 289 3.2 BTEX distribution between gaseous and aqueous phases

290 The exsufflation time required for the equilibrium to be achieved was different for rainwater samples  
 291 and ultrapure water, which suggests that physico-chemical properties and BTEX content distributed  
 292 between different phases have certain impact on the adsorption to the aqueous phase (Figure 2,  
 293 left). However, the correlations were registered only between the concentrations of B, and  $\text{F}^-$  (-0.72)  
 294 and  $\text{NH}_4^+$  (0.83) (Figure S2, Supplementary material). The impact of rainwater physico-chemical  
 295 properties on the BTEX retention was further examined by insufflating 2 ppbV of BTEX in 10 pre-  
 296 exsufflated rainwater samples, as described in our previous paper (Šoštarić et al., 2015). The results  
 297 showed slightly longer exsufflation periods for rainwater samples compared to the pure water,



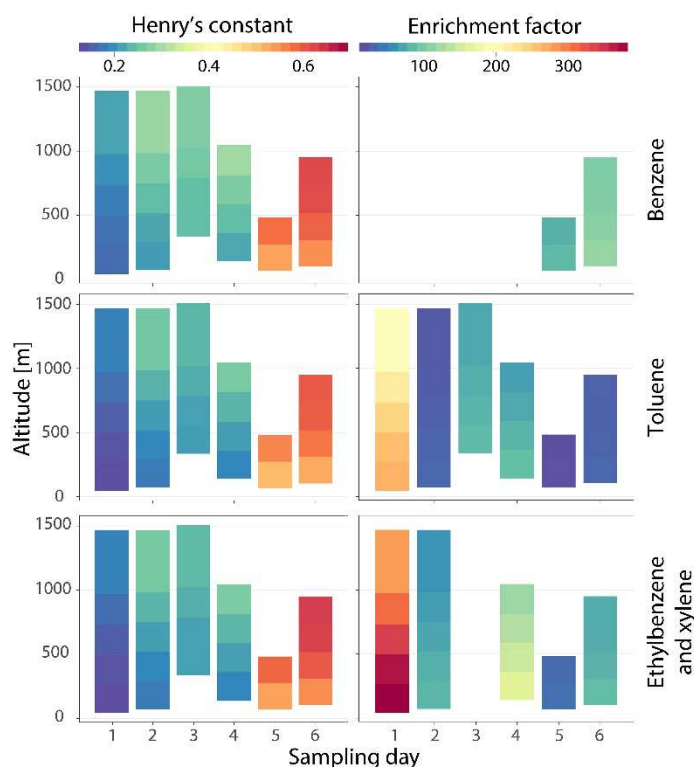
298 indicating that physico-chemical properties are not the main contributor to the extended retention in  
 299 the rainwater. The comparative qualitative analysis of rainwater and ultrapure water exsufflation  
 300 time series obtained by real-time PTR-MS measurements showed that different capacities for BTEX  
 301 retention can be mainly associated with BTEX aerosol fraction (Figure 2, right).



302  
 303 Figure 2. The time required to reach equilibrium ( $t_{eq}$ ) and 99, 95 and 90% quantity of benzene and  
 304 toluene to be exsufflated (left), and toluene exsufflation from ultra-pure water and rainwater (right).

305 Generally, due to very small  $K_H$  values of aromatic compounds, BTEX concentrations in rainwater are  
 306 expected to be low (Słomińska et al., 2014). However, according to the results, EF values were in the  
 307 range from 61 to 128, from 8 to 209, and from 25 to 295 for B, T and EX, respectively, indicating that  
 308 the BTEX amounts in the aqueous phase significantly exceeded the theoretically predicted levels.  
 309 Both lower and higher EF values have been reported in the literature. According to the studies  
 310 dealing with the distribution of chlorinated hydrocarbons and monocyclic aromatic hydrocarbons  
 311 between air and rainwater, EF ranged from 2.4 to 34 for BTEX (Okochi et al., 2004; Sato et al., 2006),  
 312 whereas the study of Valsaraj et al. (1993) reported several hundred to a thousand-fold enrichment  
 313 of hydrophobic organic compounds in fog samples. The enhanced BTEX transfer to urban dew water  
 314 was also shown by Okochi et al. (2005), with the reported EF values ranging from 7.87 to 20.2.  
 315 Furthermore, Fries et al. (2007) showed that the concentrations of aromatic hydrocarbons including  
 316 ethylbenzene, xylenes and 1,2,4-trimethylbenzene, have been significantly lower in rain ( $15\text{--}53\text{ ng L}^{-1}$ )  
 317 than in snow ( $71\text{--}2200\text{ ng L}^{-1}$ ). In the later study, Fries et al. (2008) found that in-cloud scavenging  
 318 could be a possible explanation for the occurrence of VOC in fallen snow.

319 Figure 3 illustrates the  $K_H$  and EF altitude distribution for BTEX. It should be mentioned that  $K_H$  and EF  
 320 for B and EX exhibit the similar pattern. As can be seen,  $K_H$  and consequently EF, change as the  
 321 raindrop falls to the ground.  $K_H$  values calculated using the average temperature on the path of the  
 322 raindrop differ  $\pm 20\%$  from those obtained using the rainwater temperature. Such agreement  
 323 indicates that the rainwater temperature measured at the ground level is a good indicator of  
 324 atmospheric conditions under which reactions with BTEX takes place. Moreover, in the study of Lin et  
 325 al. (2011), it was concluded that, at higher altitudes in locations with dominant local sources, VOC  
 326 concentrations were generally lower, and hence, higher  $K_H$  values would not be expected to affect  $C_{R}$ ,  
 327 nor calculated values of  $D_{OBS}$  and EF.



328

329 Figure 3. Henry's constant and EF BTEX altitude distributions.

330 As suggested by the field studies (Valsaraj et al., 1993; Goss, 1994; Okochi et al., 2004; Starokozhev  
 331 et al., 2009), as well as the laboratory experiments (Bruant and Conklin, 2000; Bruant and Conklin,  
 332 2002; Raja et al., 2002; Raja and Valsaraj, 2004; Šoštarić et al., 2016), the interfacial adsorption might  
 333 be the major mechanism associated with the enhanced VOC transfer to the aqueous phase.

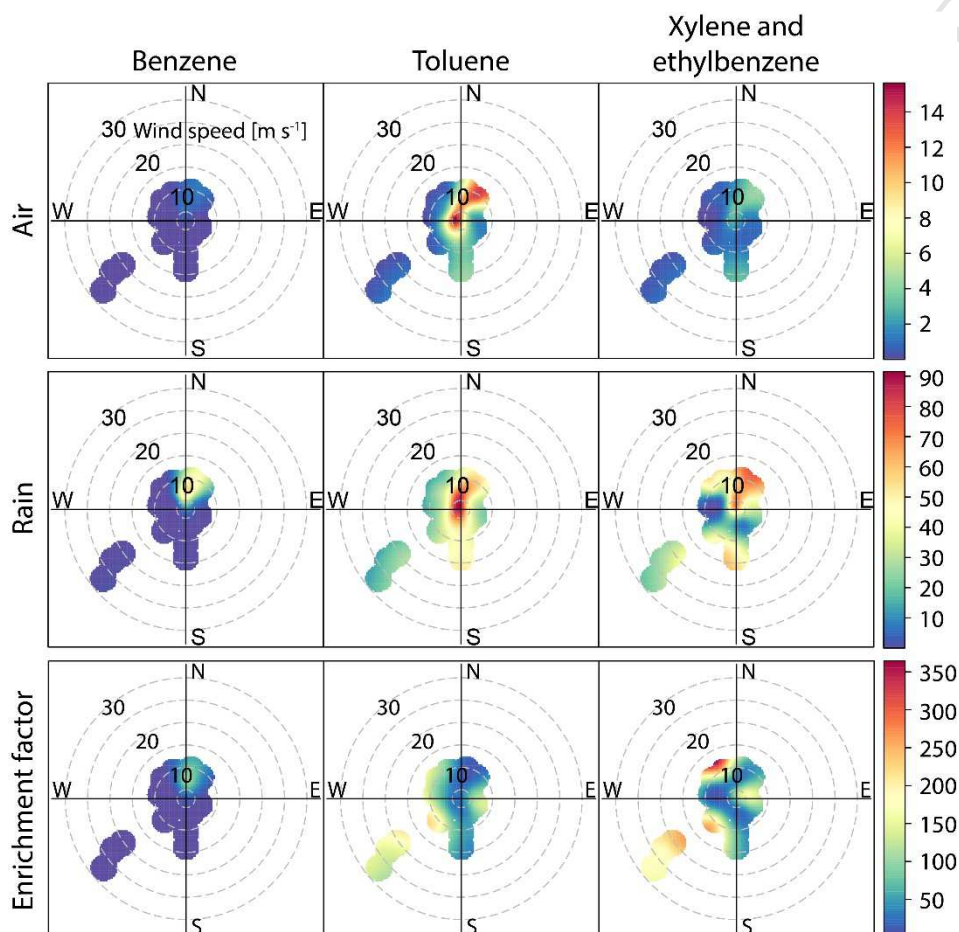
334 Some previous studies have examined the composition of atmospheric water and the impact of  
 335 different species, including nitric acid, anionic and nonionic surfactants, as well as the impact of  
 336 salinity and pH on air-water VOC distribution (O'Sullivan et al., 1996; Vane and Giroux, 2000; Sato et  
 337 al., 2006; Allou et al., 2011). The supersaturation of VOC in rain samples was explained by decreased  
 338 rainwater polarity associated with the presence of different organic compounds (Sato et al., 2006).  
 339 The presence of colloidal organic matter with its large binding capacity for many hydrophobic species  
 340 was found in fog droplets (Valsaraj et al., 1993). Similarly, in the study of Okochi et al. (2005), the  
 341 enhanced dissolution of VOC species in urban dew compared to rainwater was explained by the fact  
 342 that dew forms near the ground and contain more humic-like substances that could lead to a  
 343 decrease in water surface tension and consequently result in higher VOC enrichment.

344 The present study considers several factors that could contribute to BTEX enrichment in rainwater,  
 345 including BTEX concentrations, rainwater physico-chemical properties, rainfall intensity, air masses  
 346 origin, meteorological conditions and adsorption at the air/water interface.

347 As regards the physico-chemical characteristics of rainwater, only  $F^-$  and  $SO_4^{2-}$  can be considered  
 348 important for the prediction of  $EF_B$  (-0.73) and  $EF_{EX}$  (0.62), respectively. The latter indicates that  
 349  $SO_2$ -rich coal burning emissions are also a significant source of EX, while the results for B should be  
 350 taken with caution due to the small data size.

351 Higher rainwater T enrichment was mostly observed for low gaseous T concentrations, high TOC  
 352 ( $8-12 \text{ mg L}^{-1}$ ) and turbidity (Supplementary file 1), although the strict link between rainwater  
 353 enrichment and gaseous concentrations cannot be established. The exsufflation dynamics (Figure 2)  
 354 and the EF values suggest that prolonged BTEX retention could also be attributed to the adsorption  
 355 to aerosol fraction. We have considered T partitioning, not only because T concentrations were in a  
 356 relatively broad range, but also because of the significant number of samples collected in both

357 seasons with comparable concentrations, despite the fact that similar findings were observed for the  
 358 rainwater enrichment with EX (Supplementary file 2). As can be observed in Figure 4 and Table S6 of  
 359 the Supplementary material, higher T enrichment was mostly associated with a higher wind speed at  
 360 the sampling site (up to  $30 \text{ m s}^{-1}$ ) and air masses coming from SW area, whereas the lowest rainwater  
 361 enrichment was registered under relatively stable atmospheric conditions ( $w_s < 5 \text{ m s}^{-1}$ ). Similar  
 362 associations were also observed for EX. Higher rainwater enrichment could be the result of the  
 363 prolonged contact time between the aqueous and the gaseous phases, when strong wind-driven  
 364 raindrops were falling obliquely.



365  
 366 Figure 4. The relationship between BTEX air mixing ratios (ppb), rain concentrations (nM) and  
 367 enrichment factor and wind characteristics.

368 According to GRRF results (Table S7, Supplementary material), physico-chemical rainwater properties  
 369 and gaseous T concentrations appear to be of greater importance than meteorological factors for  
 370 predicting T and EX rainwater enrichment. Furthermore, these findings also indicate that ground  
 371 level gaseous concentrations have higher impact on the transfer of species to the aqueous phase  
 372 than the polluted air masses coming from greater atmospheric heights.

373 Out of 24 examined MVA regression methods, some of which were previously successfully applied  
 374 for prediction of  $\text{PM}_{10}$  and VOC emissions (Stojić et al., 2015d; Perišić, 2016), it has been shown that  
 375 RF, IBk and IBkLG can provide predictions of  $\text{EF}_T$  and  $\text{EF}_{EX}$  based on the variables of the highest  
 376 importance derived by GRRF with relative errors of approx. 20%, *i.e.* 27%, and correlation coefficients  
 377 around 0.95 and 0.87, respectively (Table 1). Conversely, the prediction of  $\text{EF}_T$  and  $\text{EF}_{EX}$  based on  
 378 Unmix derived source contributions was less accurate ( $K^*$ : 36.3% relative error and correlation  
 379 coefficient 0.79). As can be concluded, functional description of  $\text{EF}_T$  and  $\text{EF}_{EX}$  can be based on certain  
 380 rainwater properties and gaseous T concentrations. In addition to ambient and rainwater B

381 concentrations,  $EF_B$  is affected by meteorological conditions (sample and ambient temperature and  
382 Rh), but these results should be taken with caution due to the small B data size.

383

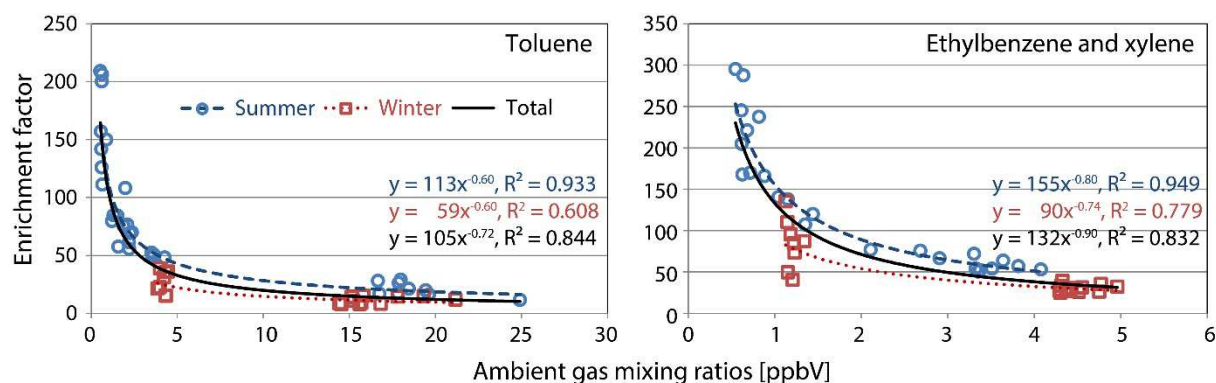
384 Table 1. MVA method performance comparison for enrichment factor prediction based on the  
385 measured parameters and Unmix-derived source contributions: absolute error, relative error and  
386 correlation coefficient (r).

Method	$EF_T$						$EF_{EX}$					
	Measured			Unmix derived			Measured			Unmix derived		
	Abs.	Rel.	r	Abs.	Rel.	r	Abs.	Rel.	r	Abs.	Rel.	r
Alternating Model Tree	18.5	34.8	0.88	40.8	76.4	0.67	33.8	42.0	0.78	59.2	111.0	0.53
Conjunctive Rule	26.1	48.9	0.64	32.0	59.9	0.40	39.2	48.7	0.55	43.4	81.2	0.41
Decision Stump	25.5	47.8	0.57	29.3	54.9	0.46	36.4	45.2	0.57	39.4	73.8	0.45
Decision Table	17.8	33.3	0.85	25.3	47.4	0.68	33.9	42.1	0.71	44.8	83.9	0.46
Elastic Net	15.2	28.5	0.88	22.3	41.9	0.77	29.4	36.5	0.74	38.6	72.4	0.62
Gaussian	16.1	30.2	0.88	21.9	41.0	0.79	30.6	38.0	0.72	41.0	76.8	0.60
IBk	11.6	21.7	0.95	20.6	38.5	0.79	22.4	27.8	0.89	31.9	59.8	0.68
IBkLG	11.6	21.7	0.95	20.6	38.5	0.79	22.4	27.8	0.89	31.9	59.8	0.68
Isotonic	13.7	25.7	0.93	31.4	58.9	0.54	34.9	43.3	0.62	39.5	74.0	0.58
K*	11.9	22.3	0.92	19.4	36.3	0.79	22.0	27.4	0.87	31.0	58.1	0.68
Least MedSq	17.4	32.7	0.87	35.6	66.8	0.67	52.7	65.4	0.29	47.4	88.8	0.50
Linear Regression	15.5	29.1	0.89	22.7	42.5	0.78	30.8	38.2	0.74	39.0	73.2	0.61
LWL	15.9	29.7	0.91	26.7	50.0	0.72	35.5	44.1	0.70	39.4	73.9	0.53
M5P	13.9	26.1	0.91	21.6	40.5	0.78	29.7	36.9	0.74	40.4	75.7	0.60
M5Rules	15.0	28.1	0.90	22.9	42.9	0.74	30.8	38.3	0.74	42.2	79.1	0.59
Multilayer	17.7	33.1	0.92	24.2	45.4	0.80	37.5	46.5	0.74	49.5	92.7	0.55
Pace Regression	15.9	29.8	0.88	22.7	42.6	0.77	30.1	37.4	0.74	38.6	72.4	0.62
Random Forest	11.6	21.7	0.95	22.5	42.2	0.76	27.8	34.5	0.76	34.9	65.4	0.62
Random Tree	13.8	25.9	0.89	28.6	53.7	0.59	32.8	40.8	0.67	41.7	78.1	0.54
RBF Network	29.9	56.0	0.45	30.4	57.0	0.65	44.8	55.7	0.49	52.8	98.9	0.40
RBF Regressor	14.7	27.6	0.91	23.6	44.2	0.78	28.1	34.9	0.83	40.9	76.6	0.56
REP Tree	13.9	26.1	0.91	22.5	42.1	0.72	36.9	45.8	0.58	42.0	78.8	0.43
Simple Linear Regression	29.9	56.1	0.49	30.3	56.7	0.55	34.3	42.7	0.71	40.5	75.8	0.61
SMOreg	14.7	27.5	0.88	21.5	40.3	0.79	28.8	35.8	0.75	39.0	73.2	0.61

387

388 Figure 5 represents EF for T and EX as a function of their ambient gas phase mixing ratios. As can be  
389 seen, EF values increased in summer, due to ambient air temperature, which is one of the most  
390 important factors for the decrease of the surface tension leading to enhanced interfacial adsorption  
391 (Bruant and Conklin 2000, 2002). Another important feature of Figure 5 is the power functional  
392 dependence of EF on ambient gas phase mixing ratios, as adsorption processes are generally more  
393 efficient for lower adsorbate concentrations and are characterized by the power functions. These  
394 findings are in compliance with the findings of Sato et al. (2006) who showed that rainwater  
395 enrichment is especially significant for the species with lower atmospheric concentrations.

396



397  
398 Figure 5. The relationship between T and EX enrichment factors and their gaseous concentrations.  
399

#### 400 4. Conclusions

401 The transfer of BTEX and other VOC from the atmosphere to various forms of atmospheric water is  
402 an important process that affects the global transport of air pollutants, environmental fate and  
403 enables the transfer of these species to terrestrial and aquatic systems. The purpose of this study  
404 was to investigate the scavenging potential of rainwater and consider the potential mechanisms and  
405 factors associated with this phenomenon.

406 As shown, BTEX concentrations observed in the aqueous phase exceeded the theoretically predicted  
407 values. Given that the interfacial adsorption is assumed to be the major mechanism underlying the  
408 enhanced rain scavenging of BTEX, the removal process was observed to be more efficient for lower  
409 gas mixing ratios, mainly due to equal surface available for smaller number of molecules and the  
410 prolonged contact time between the two phases when wind-driven rain drops were falling obliquely.  
411 Accordingly, theoretical predictions are probably more accurate in the area of larger gaseous  
412 concentrations, whereas in the case of lower concentrations, transfer to the aqueous phase is often  
413 underestimated. Furthermore, the results of the presented regression multivariate analysis suggest  
414 that multiple factors determine the spatio-temporal BTEX distribution in the environmental  
415 multiphase system, including ambient mixing ratios, physico-chemical properties of rainwater and  
416 meteorological data. More specifically, the functional description of  $EF_T$  and  $EF_{EX}$  can be based on  
417 certain rainwater properties. On the other hand, it has been shown that  $EF_B$  is affected by  
418 meteorological conditions (sample and ambient temperature, and  $R_h$ ), as well as B ambient and  
419 rainwater concentrations (however, these results should be interpreted with caution due to the small  
420 B data size).

421

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426

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**Highlights**

- Potential of rainwater for BTEX scavenging from ambient air was examined
- BTEX concentrations in rain samples exceeded the theoretically predicted values
- BTEX retention could be associated with BTEX aerosol fraction
- Random forest and instance based algorithms provide reliable enrichment predictions
- Gas mixing ratios, rainwater characteristics and meteorology affect BTEX distribution